

Water storage in the refractory lithospheric mantle

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Supplementary Information

The Supplementary Information includes:

- Methods
- Tables S-1 to S-4
- Figure S-1
- Supplementary Information References
- Additional References for Data Sources in Figures 2 and 3

Methods

Major element analysis

Major element compositions of olivine were measured using the JEOL JXA8100 electron probe microanalyzer at the Institute of Geology and Geophysics, Chinese Academy of Sciences. The operating accelerating voltage was 15 kV and beam current was 20 nA with a focused beam (spot diameter ~ 5 μm). The natural minerals and synthetic oxides that were used calibration are as following: albite (Na and Al), bustamite (Mn), diopside (Ca, Si, and Mg), K-feldspar (K), rutile (Ti), Fe_2O_3 (Fe), NiO (Ni) and Cr_2O_3 (Cr). Two TAP analytical crystals were used to analysis Na, Mg, Al, and Si ($K\alpha$), one LIFH crystal was used to analysis Cr, Mn, Fe, and Ni ($K\alpha$), and one PETJ crystal was used to analysis K, Ca, and Ti ($K\alpha$). Sodium and K were first measured in order to minimize the loss of volatiles by electron beam irradiation. The spectral line adopted of each element is Na (129.057), Mn (146.180), K (120.399), Mg (107.574), Si (76.815), Fe (134.638), Al (90.119), Ca (108.276), Ti (89.039), Cr (159.315), and Ni (115.218), respectively. The peak counting time for each element was 10-20 s. The detection limits are (1σ) 0.02 wt. % for Mg, Al, Si, Cr, K, Mn, Ti, Ni, and Fe, 0.03 wt. % for Na and Ca. A program based on the ZAF procedure was used for data correction. Detailed EPMA analytical conditions and results are shown in Table S-2.

SIMS analysis of water

Water contents in the olivine were measured using a CAMECA IMS 1280-HR secondary ion mass spectrometry (SIMS) at the State Key Laboratory of Isotope Geochemistry, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, following the method outlined by Zhang *et al.* (2020). Prior to analysis, samples were placed on double-

sided adhesive tape and enclosed in a tin-based alloy. This alloy mount was then photographed under microscope (reflected light) and then gold-coated (~30 nm thick). This innovative sample preparation method enhances the vacuum condition within the sample chamber, yielding a smooth surface that improves detection limits and precision (Cui *et al.*, 2024).

When SIMS was maintained at the optimum condition after thorough baking, an automatic liquid nitrogen refilling system was used to maintain ultra-high vacuum in the sample chamber. The sample mount was loaded into the storage chamber overnight and further pumped down (for 1 to 2 h) before the analysis (Zhang *et al.*, 2018). A Cs⁺ primary beam of 4-5 nA and $\Phi \sim 15 \mu\text{m}$ (10 kV impact energy) was used to sputter secondary ion from the samples. The size of the analytical area was $30 \mu\text{m} \times 30 \mu\text{m}$ (15 μm spot size + 15 μm rastering). The pre-sputtering process lasted for 170 seconds with a raster area of $50 \mu\text{m} \times 50 \mu\text{m}$, which is larger than the analysis area to minimize the water background signal. The normal-incidence electron gun and the nuclear magnetic resonance (NMR) controller were used to ensure charge compensation throughout the measurement session and to stabilize the magnetic field, respectively. Other instrument parameters include 2000 μm field aperture, 400 μm contrast aperture, 60 μm entrance slit, 80 μm max area, 173 μm exit slit, and 50 eV energy slit with 5 eV gap. The Faraday Cup detectors, located on L'2 trolley with 10^{10} Ohm amplifiers, and an electron multiplier of mono-collector were used to measure the $^{16}\text{O}^-$ and $^{16}\text{O}^1\text{H}^-$, respectively. For $^{16}\text{O}^-$, 500 mm collector slits were used to yield a ~2500 mass resolution power (MRP) with sufficiently flat plateau, while a ~173 μm collector slit was used for electron multiplier (yielding ~7000 MRP) to obtain a sufficiently flat plateau and separate isobaric interference of $^{17}\text{O}^-$ on $^{16}\text{O}^1\text{H}^-$. Background H₂O content was determined on standard San Carlos olivine as it has the lowest H₂O abundance among all standards. Water content was calculated according to the calibration curves established by analyses of the olivine standards of San Carlos (Fo = 91.1), KLB-1 (Fo = 89.3), ICH-30 (Fo = 90.6), and Mongok (Fo = 99.0) (Zhang *et al.*, 2020). The uncertainty of the calibrated water content is ~10 %. All data measured on standards and samples are listed in Table S-3 and S-4, respectively. The plot with H₂O calibration on the standards is shown in Table S-3.

Supplementary Tables

Table S-1 Water contents ($\mu\text{g.g}^{-1}$) and major oxide compositions (wt. %) of olivine from Bursa and Lycian ophiolites in Türkiye.

Table S-2 Results of standards and working conditions of EPMA analysis.

Table S-3 Results of standards and the plot of water calibration on the standards of the SIMS analysis.

Table S-4 All SIMS data of samples.

Tables S-1 to S-4 are available for download (.xlsx) from the online version of this article at <http://doi.org/10.7185/geochemlet.2530>

Supplementary Figures

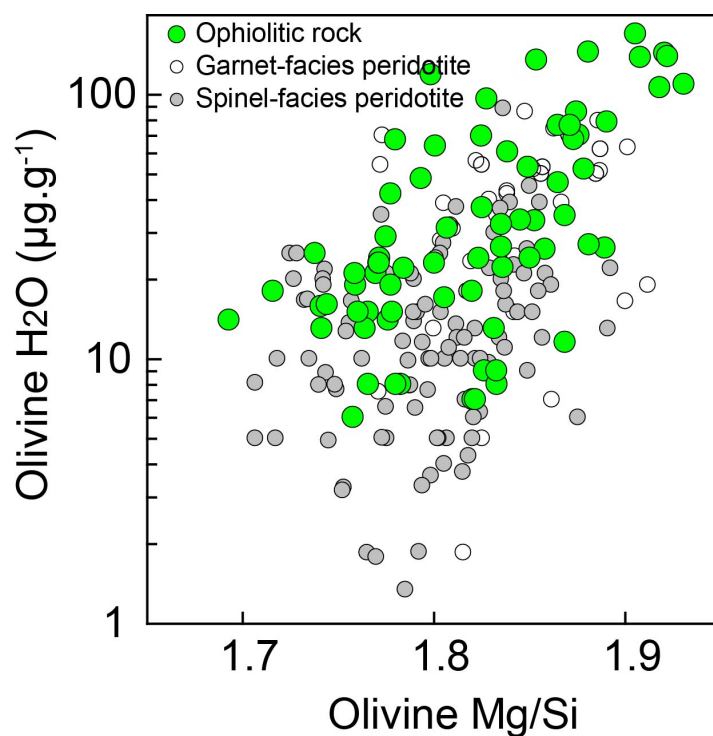


Figure S-1 Plot of H₂O (μg.g⁻¹) vs. Mg/Si for olivine from ophiolites, and mantle xenoliths. Data sources are as in Fig. 3.

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